## Modeling atmospheric and ocean dynamics in the lab



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The test of all knowledge is experiment. Experiment is the sole judge of scientific truth.

— Richard P. Feynman —

## The minimal model of the weather system/ ocean circulation (mid-latitudes)

Rotation +
meridional temperature
difference = weather
(or ocean circulation)

• Idea: Let's construct minimalistic laboratory analogs to better understand the basic underlying dynamics



• A differentially heated cylindrical tank, mounted on a turntable. "Rotating annulus"

#### **Dimensions (Cottbus):**

- a = 45 mm
- b = 120 mm
- d = 135 mm



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#### **Dimensions (Budapest):**

- a = 45 mm
- $b = 150 \, mm$
- d = 40 mm



• A differentially heated cylindrical tank, mounted on a turntable. "Rotating annulus"

#### **Dimensions (Tallahassee):**

- a = 160 mm
- $b = 610 \ mm$
- d = 80 mm







"Sideways convection" – no threshold in ΔT (i.e. No 'critical Rayleigh number')

*Any* temperature difference can initiate the flow



cooling

heating

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cooling

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cooling

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**Rotation!** 



#### Baroclinic instability

![](_page_12_Picture_1.jpeg)

cooling

heating

![](_page_12_Picture_4.jpeg)

Zonal flow (thermal wind) Geostrophic theory: Tilted density surfaces

![](_page_12_Picture_6.jpeg)

 $-2\Omega \vec{e_z} \times \vec{u}$ 

#### Baroclinic instability

![](_page_13_Picture_1.jpeg)

![](_page_13_Picture_2.jpeg)

**KOTATION** Zonal flow

(thermal wind) Geostrophic theory: Tilted density surfaces

![](_page_13_Picture_5.jpeg)

#### Baroclinic instability

![](_page_14_Picture_1.jpeg)

![](_page_14_Picture_2.jpeg)

(thermal wind) Geostrophic theory: Tilted density surfaces

![](_page_14_Picture_4.jpeg)

# Baroclinic waves

## control parameters:

- rotation rate, radial temperature difference
- Different planetary atmospheres can be modelled
- <u>Venus:</u> slow rotation, zonal flow
- <u>Earth:</u> fast rotation → Coriolis effect, cyclones ("weather")

![](_page_15_Picture_6.jpeg)

![](_page_15_Picture_7.jpeg)

![](_page_15_Picture_8.jpeg)

![](_page_15_Picture_9.jpeg)

# Baroclinic waves

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![](_page_16_Picture_6.jpeg)

![](_page_16_Picture_7.jpeg)

![](_page_16_Picture_8.jpeg)

![](_page_16_Figure_9.jpeg)

![](_page_17_Picture_0.jpeg)

The regime diagram (after Fultz)

![](_page_18_Figure_1.jpeg)

- Large annulus → very small Rossby numbers (quasigeostrophic turbulence)
- Temperature time series from fixed locations

• Question: How do the statistical properties of extreme events in this model weather (and climate) depend on the meridional temperature contrast?

• The distribution of fluctuations widens with the increasing temperature contrast  $\rightarrow$  increasing Rossby number (~  $\Delta$ T).

![](_page_20_Figure_2.jpeg)

• The distribution of fluctuations widens with the increasing temperature contrast  $\rightarrow$  increasing Rossby number (~  $\Delta$ T).

![](_page_21_Figure_2.jpeg)

Taking the histograms of the (smoothed)

time derivatives

of the temperature fluctuation time series, and their "width" based on the Q0.05 and Q0.95 quantiles

→ data collapse via *advective scaling* (dT/dt ~  $U \cdot \nabla T \sim \Delta T^2/(\Omega L)$ )

![](_page_22_Figure_5.jpeg)

**Question:** How do the statistical properties of extreme events in weather and climate depend on the meridional temperature contrast?

## Partial answers:

If only Rossby wave dynamics was involved (minimum mid-latitude atmosphere), a marked polar amplification would yield

- a narrower distribution of temperature fluctuations, whose width scales with  $\sim \Delta T$
- generally smaller "jumps" in temperature, scaling with  $\sim \Delta T^2$
- the correlation timescale ("persistence") of the weather would significantly increase, scaling with  $\sim 1/\Delta T$

### The present-day ACC: temperature field

![](_page_24_Figure_1.jpeg)

Large meridional SST gradient, accompanying the eastward geostrophic current.

![](_page_24_Picture_3.jpeg)

The "minimalistic approach" to the lateral thermal boundary conditions – differential heating (in reality the scale of the ACC jet is set by the β-effect: Rhines scale)

# Motivation: The eocene-oligocene transition (EOT) $\sim 34~{\rm Mya}$

![](_page_25_Figure_1.jpeg)

M. Rebesco et al., Marine Geology 352 (2014)

![](_page_25_Picture_3.jpeg)

#### Motivation: The eocene-oligocene transition (EOT)

• Rapid onset of glaciation of Antarctica ca. Mya (Coxall et al., 2005)

• Kennett (1977): this major cooling episode was triggered by the opening of Drake Passage. (DP opening  $\rightarrow$  ACC  $\rightarrow$  isolating Antarctica)

• This has been challenged since then: maybe long-term decrease in atmospheric CO2 level was the main driver of Antarctic glaciation (DeConto and Pollard, 2003, Pearson et al., 2009)

• Impact events! (Ivany et al., 2000)

• What do we do? Experimental minimal modeling ("toy model") of the pre-EOT overturning of the Southern hemisphere.

![](_page_26_Figure_6.jpeg)

Steerman et al., Syst. Biol. 58(6):573-585, 2009

#### "Closing the Drake Passage" with a removable radial barrier

![](_page_27_Figure_1.jpeg)

#### Surface temperature patterns (IR)

![](_page_28_Figure_1.jpeg)

Bozóki, T. et al. (2019), Deep Sea Res. II., 160, 16-24.

![](_page_28_Figure_3.jpeg)

## Scaling properties

A linear scaling between the azimuthal and radial temperature contrast for the closed experiment runs.

$$\langle \Delta T_B \rangle \sim \langle \Delta T_r \rangle$$

$$\eta(t) = \Delta T_B / \Delta T_r$$
$$\eta_0 = 0.22$$

![](_page_29_Figure_4.jpeg)

Bozóki, T. et al. (2019), Deep Sea Res. II., 160, 16-24.

### The effect of opening on the zonal spatio-temporal variability

![](_page_30_Figure_1.jpeg)

The effect of opening on the zonal spatio-temporal variability

![](_page_31_Figure_1.jpeg)

Increasing mean "SST" in the Southern ocean following the opening – an apparent contradiction with paleoclimate. Why?

Control experiments with thermometers (Cottbus tank)

Larger thermal separation between the surface and the bulk in the open configuration!

![](_page_32_Figure_3.jpeg)

Increasing mean "SST" in the Southern ocean following the opening – an apparent contradiction with paleoclimate. Why?

Control experiments with thermometers (Cottbus tank)

Larger thermal separation between the surface and the bulk in the open configuration!

![](_page_33_Figure_3.jpeg)

Increasing mean "SST" in the Southern ocean following the opening – an apparent contradiction with paleoclimate. Why? – NUMERICS!

- T21 PlaSim + LSG ("large scale geostrophic") ocean module (22 layers, resolution: 4°×4°)
- 1000 years of simulation
- Topography, orography and ocean topography modiefied (not real paleogeography, only a ca. 100 km-wide meridional "dam" inserted into the Drake passage for the "closed" runs)

4 runs (Drake passage open/closed × two different configurations):

- $CO_2$ : 750 ppm (as 34 Mya) & sea ice + land ice ON
- CO<sub>2</sub>: 750 ppm (as 34 Mya) & sea ice + land ice OFF

PlaSim (UH Met. Institute) – a "medium complexity" GCM

https://www.mi.uni-hamburg.de/en/arbeitsgruppen/theoretische-meteorologie/modelle/plasim.html

# Significant cooling in the Southern ocean – only when the sea ice module (albedo feedback) is ON!

![](_page_35_Figure_1.jpeg)

### Summary & Conclusions

- Minimalistic laboratory "toy model" of the closed Drake passage setting before and after EOT (*very* conceptual).
- In the closed case a temperature contrast can develop between the 'Atlantic' and 'Pacific' side, that is ca. <sup>1</sup>/<sub>4</sub> of the meridional temperature difference in the domain
- After opening the passage, mean SST in the set-up slightly *increases*, and the cold extremes (long left tail) disappear.
- Turns out (from PlaSim simulations) that without active seaice dynamics, the Souther Ocean's mean SST would not decrease either, pointing to the importance of ice albedo feedback in the EOT.

![](_page_37_Picture_0.jpeg)